Weakly Compressible Two-layer Shallow-Water Flows with Friction and Entrainment along Channels

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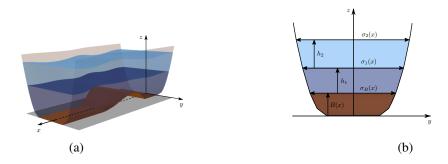


Figure 1: Schematic of a two layer channel flow with the two layers moving in opposite directions: (a) Full 3D view of the flow, (b) channel cross section.

We present a weakly compressible approach to describe two-layer shallow water flows in channels with a bottom topography and arbitrary cross sections. See the schematic in Figure 1. These flows are characterized by a large horizontal length scale relative to their depth. Applications of this model include the study of internal waves observed in rivers and possibly caused by shear flow instabilities [2]. Shallow-water flows are typically modeled by the Saint-Venant equations, a hyperbolic balance law that results from the *depth averaging* of the Euler equations.

Computations of two layer flows in channels are more challenging than those occurring in one layer. Theoretical difficulties arise due to the presence of non-conservative products. For instance, the standard approach for those flows results in a conditionally hyperbolic balance law with non-conservative products. See [5] for more details. In this work, we consider the fluid to be weakly compressible and follow a relaxation approach as in [1]. The corresponding model is unconditionally hyperbolic. See [4] for a related work in channels with vertical walls and constant width. In the talk, a detailed description of the properties of the model will provided, including entropy inequalities. Furthermore, numerical experiments using a central-upwind scheme will be presented.

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1 Details of Model

The model for the incompressible two-layer shallow water flows in channels can be found in [3]. After adding source terms due to friction and entrainment and considering the weak compressibility of both layers, it has been re-written as

(1a)
$$\frac{\partial(\rho_1 A_1)}{\partial t} + \frac{\partial(\rho_1 Q_1)}{\partial x} = \rho_1 S_e,$$

(1b)
$$\frac{\partial A_1}{\partial t} + u_1 \frac{\partial A_1}{\partial x} = \frac{\mu \left(p_1 - p_o - g \int_B^{w_1} (w_1 + rh_2 - z) \sigma(x, z) \right) dz}{\rho_1 c_1^2}$$

$$(1c) \qquad \frac{\partial (\rho_1 Q_1)}{\partial t} + \frac{\partial}{\partial x} \left(\rho_1 A_1 u_1^2 + \rho_1 p_1 \right) = g \rho_1 \left(I_1 - h_1 \sigma_B B' + r h_2 \frac{\partial A_1}{\partial x} \right) + \rho_1 S_{f,1} + \rho_1 S_e u_1, + p_o \frac{\partial \rho_1}{\partial x} + \rho_1 S_e u_2 + \rho_1 S_e u$$

(1d)
$$\frac{\partial(\rho_2 A_2)}{\partial t} + \frac{\partial(\rho_2 Q_2)}{\partial r} = -r\rho_2 S_e,$$

(1e)
$$\frac{\partial A_2}{\partial t} + u_2 \frac{\partial A_2}{\partial x} = \frac{\mu \left(p_2 - p_o - g \int_{w_1}^{w_2} (w_2 + rh_2 - z) \sigma(x, z) \right) dz}{\rho_2 c_2^2}$$

$$(1f) \qquad \qquad \frac{\partial (\rho_2 Q_2)}{\partial t} + \frac{\partial}{\partial x} \left(\rho_2 A_2 u_2^2 + \rho_2 p_2 \right) = g \rho_2 \left(I_2 - h_2 \sigma_1 \frac{\partial w_1}{\partial x} \right) + \rho_2 S_{f,2} - r \rho_2 S_e u_2 + p_o \frac{\partial \rho_2}{\partial x}.$$

Here h_1 , and h_2 denote the depth of each layer, u_1 , and u_2 the the cross-sectional velocities; g the acceleration of gravity, B(x) describes the bottom topography, and $\sigma(x,z)$ the width of the channel; $A_1=\int_B^{w_1}\sigma(x,z)\,dz$, and $A_2=\int_{w_1}^{w_2}\sigma(x,z)\,dz$ are the cross-sectional wet areas in each layer; $w_1=B+h_1$ denotes the total elevation of the internal layer and $w_2=B+h_1+h_2$ that of the external layer; and $Q_1=A_1u_1$, and $Q_2=A_2u_2$ are the flow rates or discharges for the internal and external layers respectively. Furthermore, $\sigma_B(x)=\sigma(x,B(x)), \sigma_1(x,t)=\sigma(x,w_1(x,t)), \sigma_2(x,t)=\sigma(x,w_2(x,t))$ denote the channel width at the bottom topography, and at the internal and external layers respectively. We note that σ_1 and σ_2 depend both on space and time since they are evaluated at the internal and external layers. The ratio of densities is denoted by $r=\rho_2/\rho_1\leq 1$. The relaxation parameter is given by μ . The vertically integrated hydrostatic pressures of both layers are given by

(2)
$$p_2 = p_o + c_{2o}^2(\rho_2 - \rho_2^o) + g \int_{w_1}^{w_2} (w_2 - z) \, \sigma(x, z) \, dz, \quad p_1 = p_o + c_{1o}^2(\rho_1 - \rho_1^o) + g \int_{B}^{w_1} (w_1 + rh_2 - z) \, \sigma(x, z) \, dz.$$

The source terms I_1 , and I_2 correspond to the vertically integrated pressure terms due to width variation, and are given by

(3)
$$I_1 = I_1(x,t) = \int_B^{w_1} (w_1 - z) \, \sigma_x(x,z) \, dz, \quad \text{and} \quad I_2 = I_2(x,t) = \int_{w_1}^{w_2} (w_2 - z) \, \sigma_x(x,z) \, dz.$$

The friction for each layer, and entrainment terms are given by $S_{f,1}$, $S_{f,2}$, and S_e respectively. The speed of sound for both layers are respectively given by

(4)
$$c_1 = \sqrt{g\left(\epsilon + r\frac{\sigma_1}{\sigma_2}\right)\frac{A_1}{\sigma_1}} = \sqrt{g\left(r\frac{A_1}{\sigma_2} + \epsilon\frac{A_1}{\sigma_1}\right)}, c_2 = \sqrt{\frac{gA_2}{\sigma_2}},$$

where $\epsilon = 1 - r$.

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